

IMPLICATIONS OF AN UPDATED VOLCANOLOGICAL CONCEPTUAL MODEL AT LOS HUMEROS FOR GEOTHERMAL EXPLORATION AND MODELING

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Thanks to authors of Deliverable 3.2:

Consiglio Nazionale delle Ricerche (CNR): M. Bonini, D. Maestrelli, G. Corti, D. Montanari, G. Moratti, G. Norini, G. Gola, A. Manzella E. Trumpy,

Università degli Studi di Bari Aldo Moro (UNIBA): D. Liotta

British Geological Service (BGS): C. Rochelle, A. Kilpatrick, A. Lacinska, J. Rushton, S. Kemp,

Technische Universität Darmstadt (TUDA): L. Weydt, K. Bär

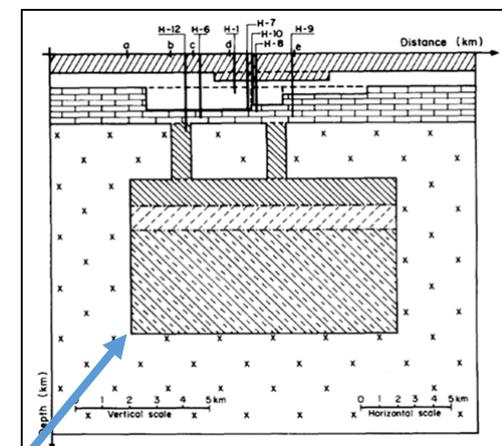


Temperature Field Simulation under Los Humeros Caldera

189

Table 1. Some physical and chemical parameters of Los Humeros volcanic center (after Verma, 1985b)

| Formation/flow | Magma erupted (km ³) | Age (My) | Temp. °C | SiO ₂ % | Minerals left behind (km ³)* |
|-----------------------------|----------------------------------|-----------|----------|--------------------|--|
| Ol. basalt | 0.25 | <0.02 | 1070 | 49 | — |
| Rhyodac.-dacites | 10 | 0.03–0.02 | 910 | 59–69 | 70 |
| <i>El Xalapazco caldera</i> | | | | | |
| Cuicuiltic tuff | 0.1 | | | 69–72 | 1 |
| Limón and other andesite | 6 | 0.04–0.02 | 960 | 56–59 | 20 |
| Llano andesite ignim. | 0.1 | | | | |
| Xoxoctic tuff | 0.6 | | 890 | 65 | 2 |
| Cueva Ahumada lavas | 0.1 | 0.06 | | | |
| <i>Los Potreros caldera</i> | | | | | |
| Zaragoza ignimbrite | 10 | 0.10 | 880–920 | 54–71 | 64 |
| Zaragoza basal fall | 2.0 | | 855 | 71 | 20 |
| ??????? | | | | | |
| Faby tuff | 10 | 0.24 | 875 | 69–73 | 100 |
| Post-Xal. rhyolites | 4.7 | 0.36–0.22 | | 73–76 | 54 |
| <i>Los Humeros caldera</i> | | | | | |
| Xaltipan ignimbrite | 115 | 0.46 | 800–875 | 69–77 | |
| Pre-Xal. rhyolites | 0.1 | 0.47 | | | |
| Teziutlán formation | 60 | 3.5–1.6 | | | |



*The parameter “Minerals left behind (km³)” refers to the volume of crystals separated and left behind in the magma chamber and is based on a fractional crystallization model of Verma (1984).

Verma et al. (1990)

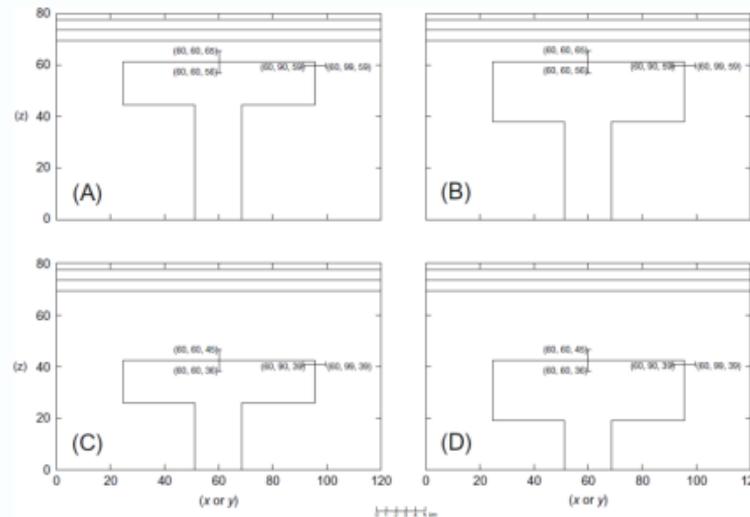
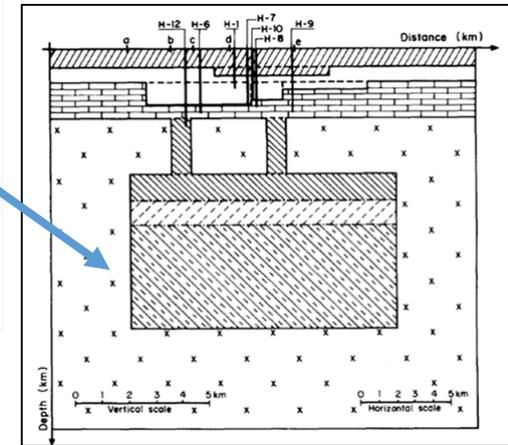


Previous volcanology-based thermal model

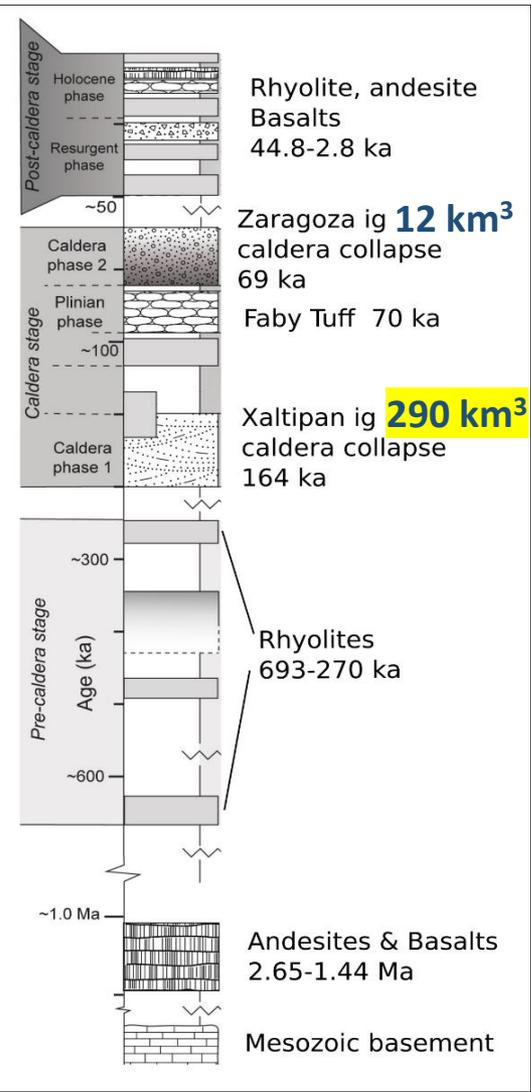
Table 2. Emplacement conditions for sensitivity evaluation of the Los Humeros geothermal field (LHC).

| Physical property (units) | Emplacement of magma chamber |
|--|------------------------------|
| Depth of the top the chamber (d_{cham}) (km) | 5–10 |
| Volume (V_{cham}) (km^3) | 1000–1400 |
| Thickness (E_{cham}) (km) | 4.5–6.5 |
| Radius (r_{cham}) (km) | 8.5 |
| Magma emplacement temperature (T_{cham}) ($^{\circ}\text{C}$) | 1350 |

Verma et al. (2011) and Verma and Gomez-Arias (2013)



Longevity and volume of the deep magmatic heat source

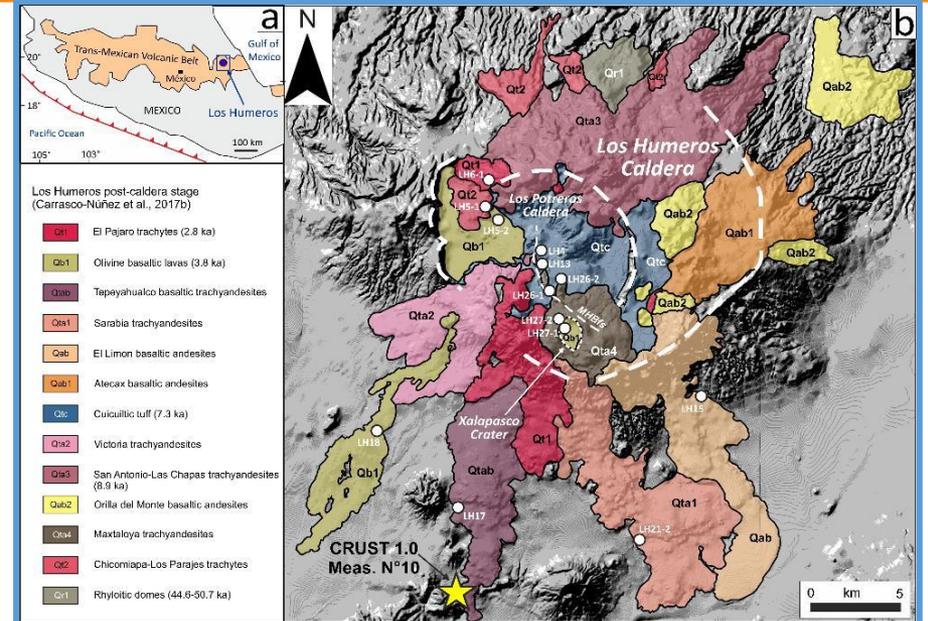


Post-peak phase (69 ka-present)
Bimodal
 magmatism/recharge

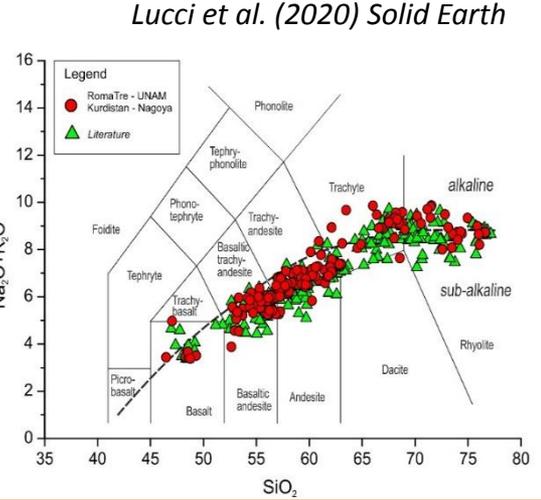
Peak phase (164-69 ka)
 Largest reservoir of rhyolitic magma and its emptying

Build up phase (>693-164 ka)
 Intrusion/Recharge /Differentiation

Carrasco-Núñez et al. (2018) G-Cubed
Cavazos & Carrasco-Núñez (2020) JVGR



RECHARGE RATES
 HAVE BEEN NOT HIGH ENOUGH TO PROMOTE LARGE INTERCONNECTION AND REHOMOGENIZATION TO A SINGLE MAGMA BODY



Petrological modeling of the LHPCS magmatic heat sources: a window into present plumbing system

LH18 (Texcal lava) is characterized by textures typical of basalts erupted directly by the deep reservoir.

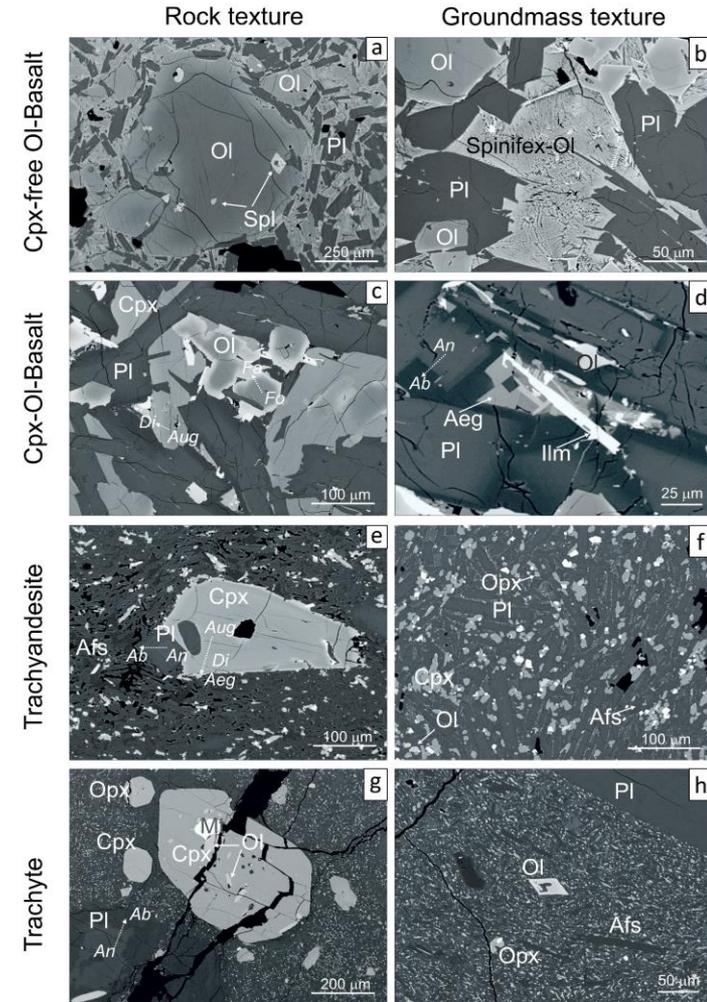
All LHPCS volcanic rocks do not show disequilibrium textures typical of assimilation and fractional crystallization (AFC) mixing processes.

The Rayleigh fractional crystallization (RFC) model

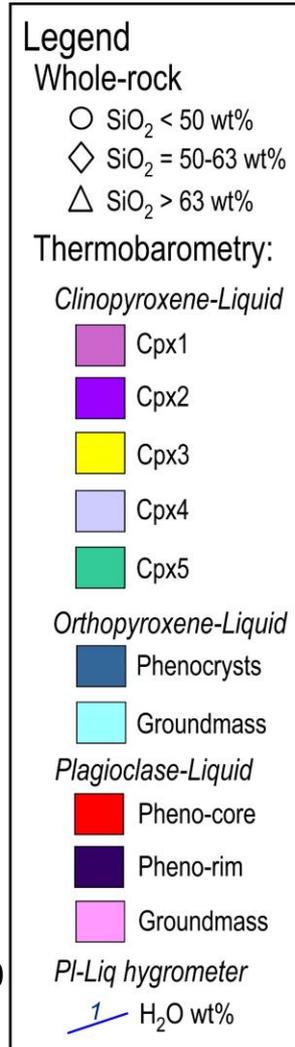
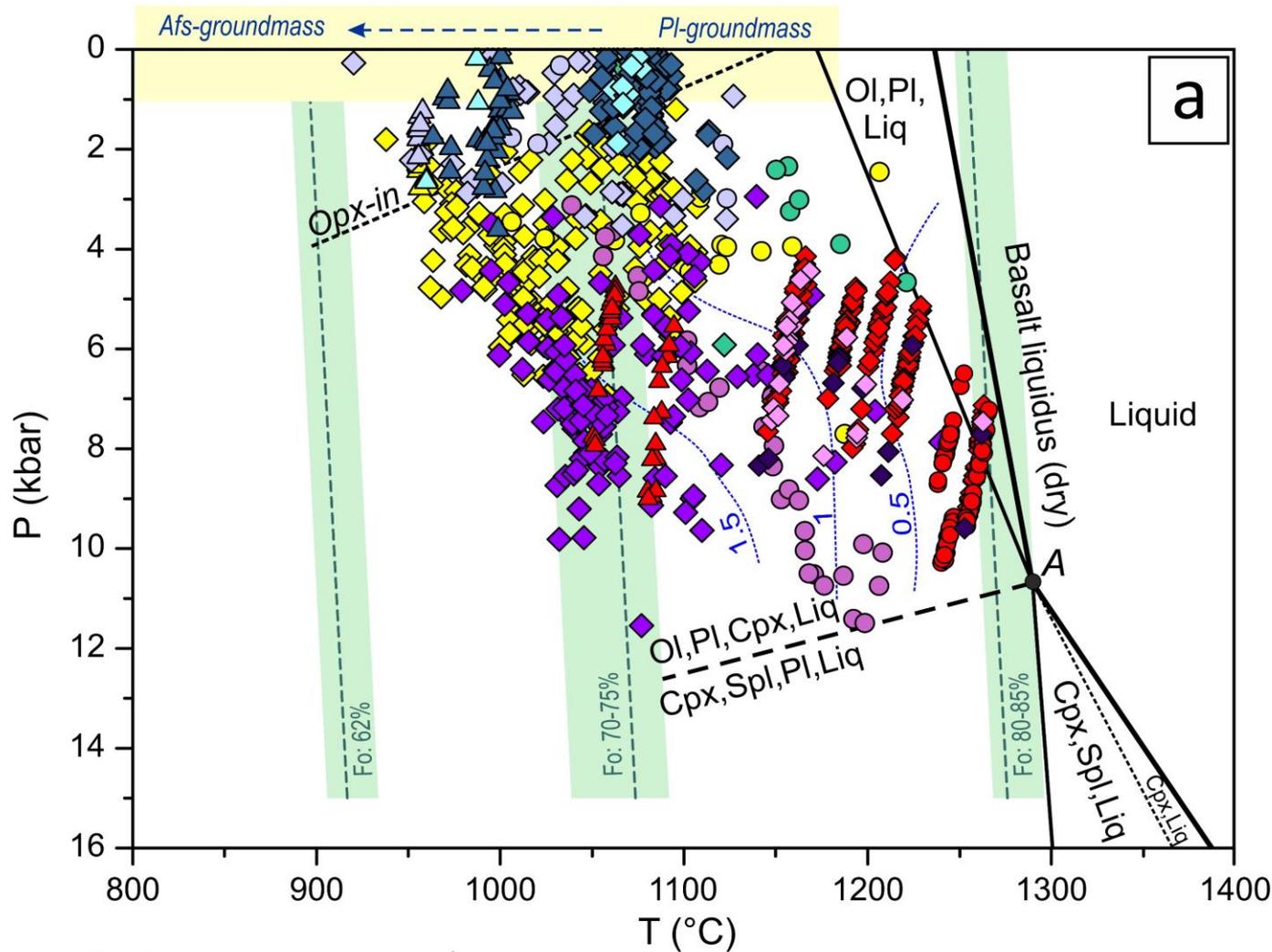
$$\sum r^2 = \sum_{i=1}^n (b'_i - b_i)^2$$

demonstrate that all LHPCS magmas, from basalts to trachytes, belong to the same line of descent and evolve through a progressive fractionation of the Pl + Cpx + Ol + Spl mineral assemblage.

Trachyandesites and trachytes represent different degrees of fractionation (RFC values in the range 45 %–74 %) starting from a Cpx-bearing basaltic source.



Geothermobarometry of the heat sources



2400 EMPA spot - 1300 min-liq pairs

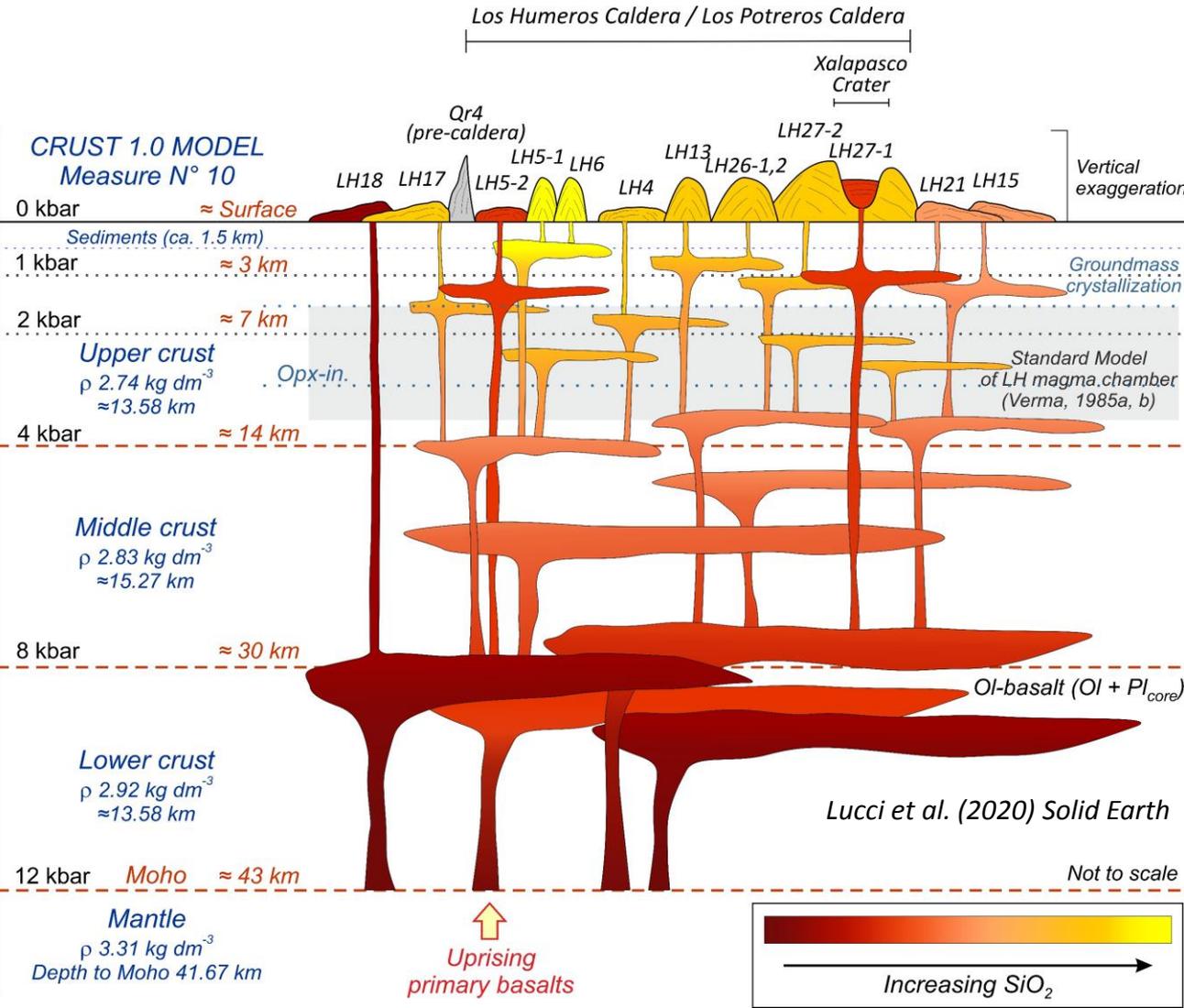
Lucci et al. (2020) *Solid Earth*



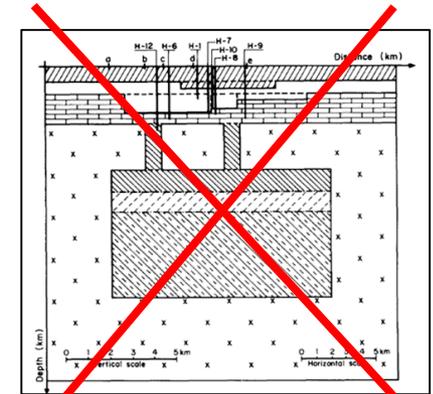
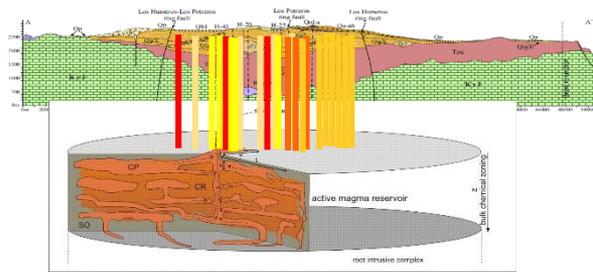
< Potsdam 2020 >
<Giordano et al.>



Step-change volcanological conceptual model of the heat sources



PRESENT Plumbing system



Volcanological conceptual model of the heat sources



| | Xaltipan | Zaragoza | LH11,12 | LH6° | LH27-2° | LH 17° (Tepeya) | LH15° (Limon) | LH5-2° (Los Potrerros) | LH18 (Texcal) |
|---|-------------------------|-----------------------|------------------------|----------|--------------------|--------------------|----------------------------|---------------------------|------------------|
| Erupted Composition | high-silica rhyolite | rhyolite- andesite | rhyolite (obsidian) | Trachyte | Trachyand esite | Trachyandesite | Opx-free Trachyandesite | Ol-Basalt | Ol-Basalt |
| Eruptive temperature*° (°C) | 790-890 | | 800 | 950 | 1050 | 1050 | 1050 | 1150 | 1250 |
| DRE Erupted Volume (km3) | 290" | 15 | 0.03 | 0.1 | 1 | 1 | 1 | 0.2 | |
| Age of eruption (ka)^ | 164 | 69 | 7 | 2.8 | >7 | >2.8 | ? | 3.8 | |
| Residence time of differentiated magma in shallow reservoir | 5ka^ | ka | a | a | a | a | a | a | a |
| Amount of evacuation of shallow reservoir | complete* | complete* | | | | | | ? | |
| Residual volume >30 km (km3) | 3600 | 161 | 0.3 | 0.2 | 1 | 1.2 | 1 | 0.02 | |
| Residual volume 30-14 km (km3) | 370 | 17 | 0.03 | 0.06 | 0.6 | 0.3 | | | |
| Residual volume 14-7 km (km3) | | | 0.05 | | | | | | |
| Residual volume <7 km (km3) | 540 | 21 | 0.004 | 0.06 | 0.4 | | | 0.2 | |
| Depth of top of shallowest transient reservoir° (km) | <=7 | <=7 | 1 [H26] | 3 | <7 | <7 | >10 | <= 3 | |
| Depth of top of shallowest fractionated/intrusive resident reservoir° (km) | <=7 | <=7 | | 3 | | | >10 | | |
| Temperature of shallowest fractionated/intrusive resident reservoir° (km) | 1000 | 1000 | 1000 | 1000 | 1050 | 1050 | 1075 | 1150 | |
| x-y area | 200 | 100 | | | | | | | |
| z-thickness | 2.7 | 0.2 | | | | | | | |
| Intrusion shape | cylinder | cylinder | sphere | sphere | sill | sill | sill | sill | sill |

*Ferriz and Mahood 1986, J Petrol

°Lucci et al 2020, Solid Earth

"Cavazos & Carrasco Nunez 2020, JVGR

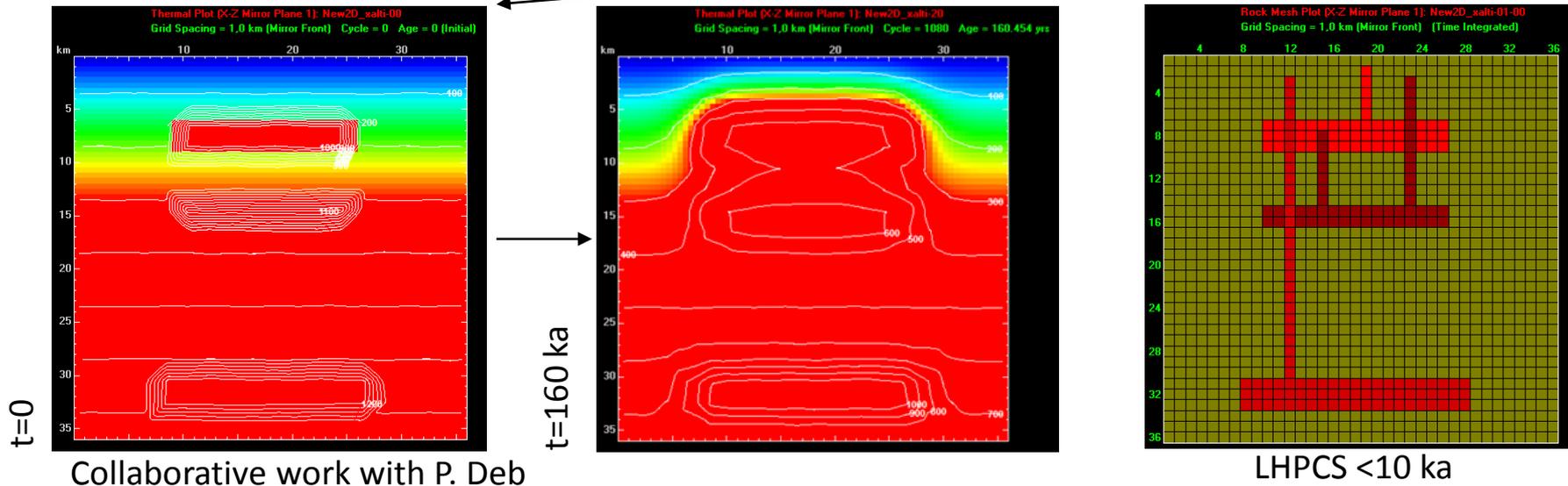
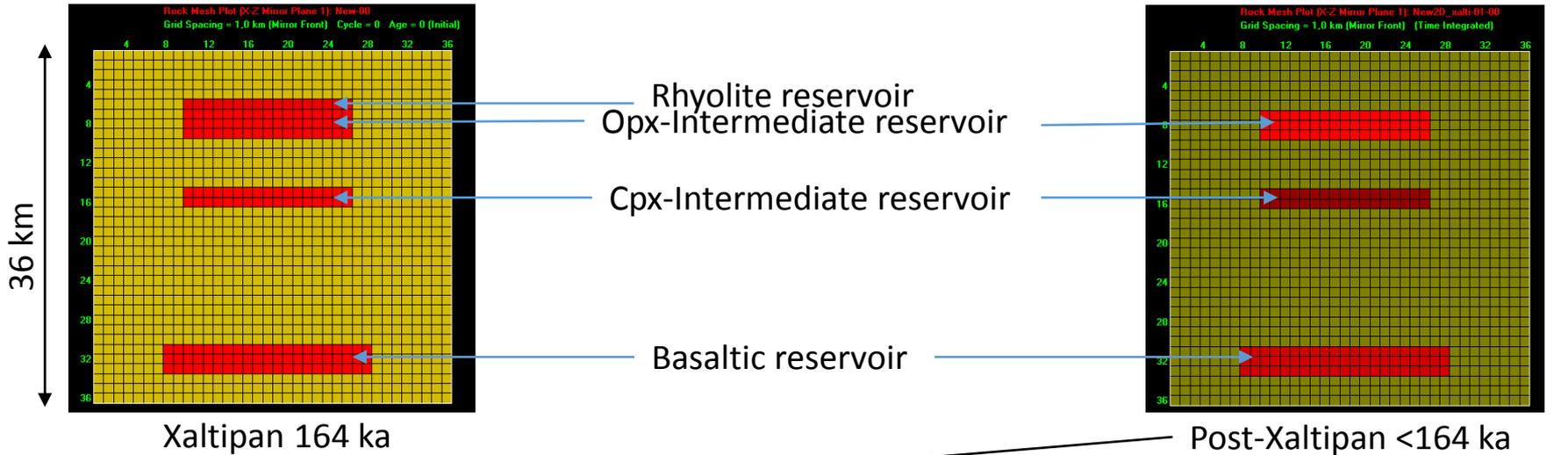
^Carrasco Nunez et al. 2018, G-Cubed



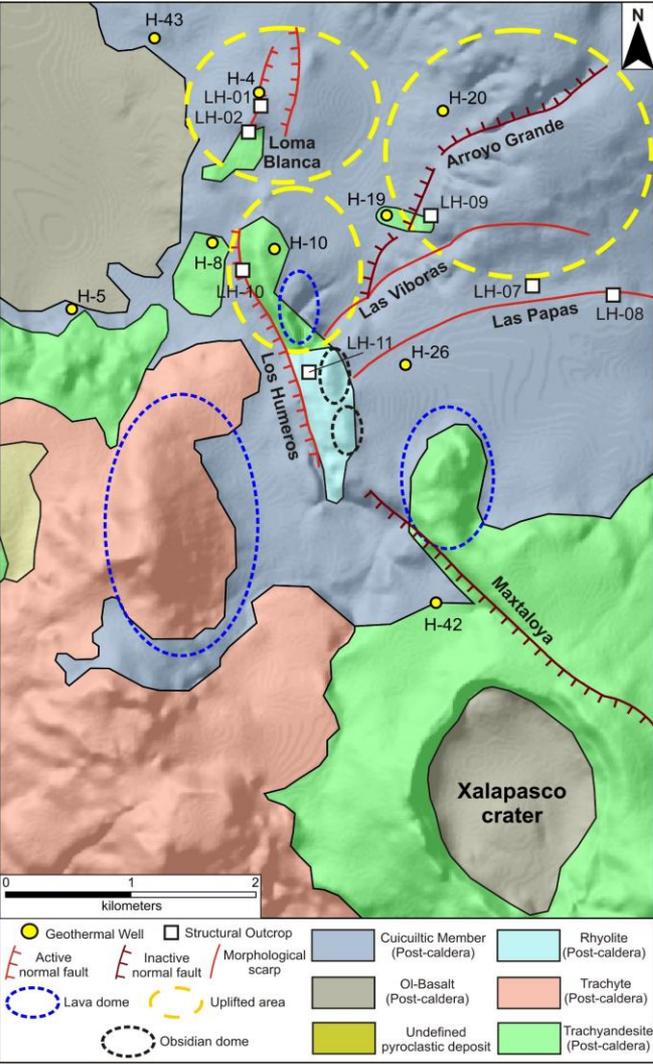
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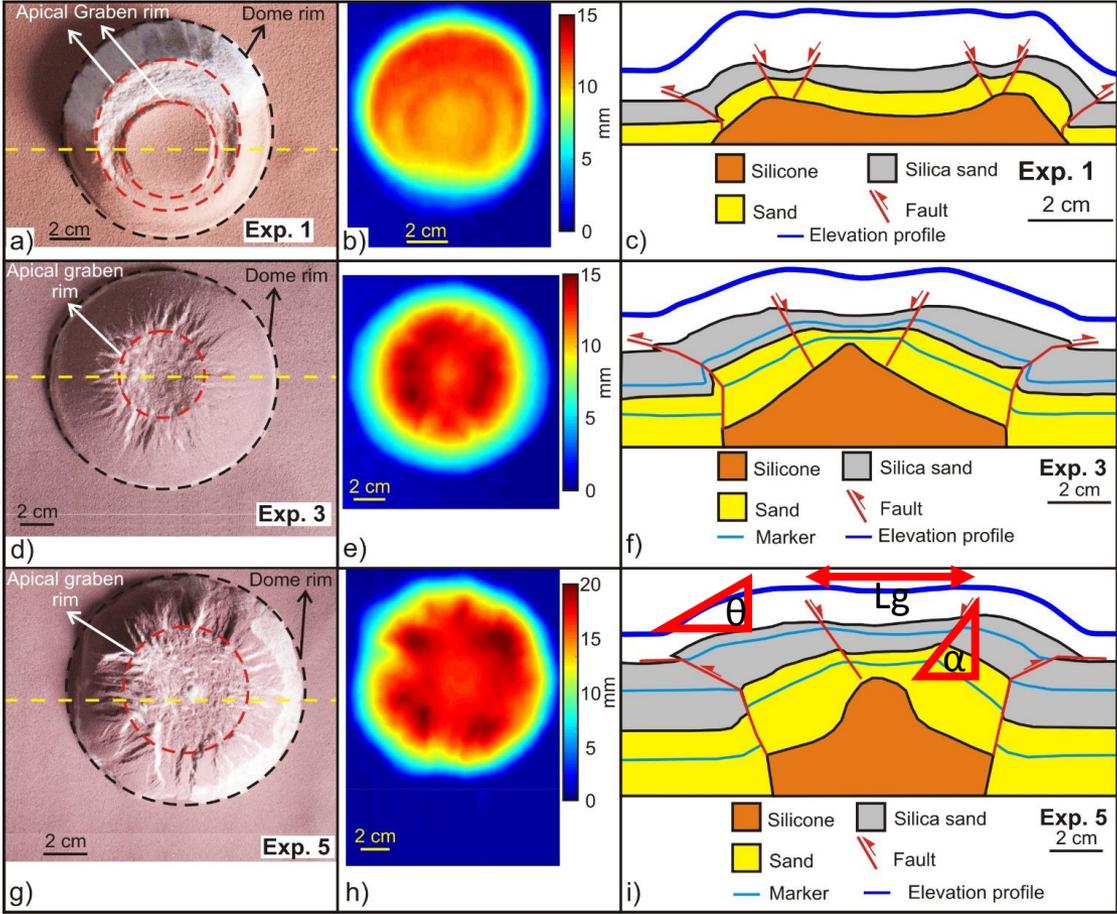
Implications for the heat sources



Implications for shallow structures



Urbani et al. (2020) Solid Earth



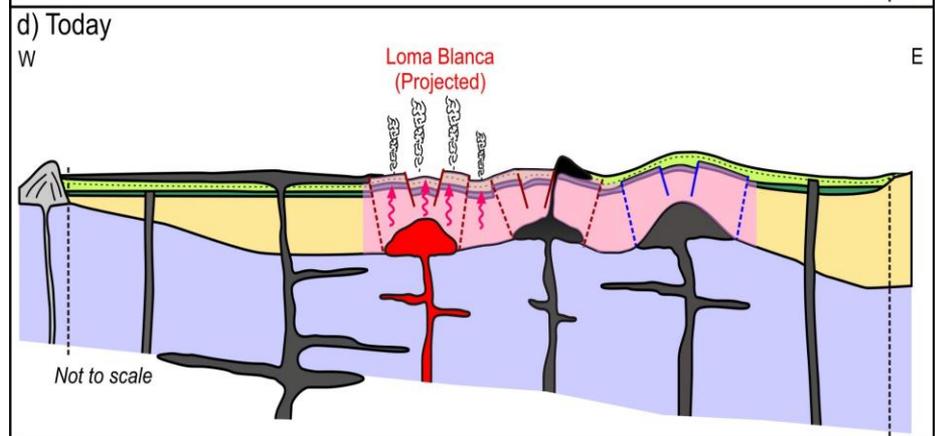
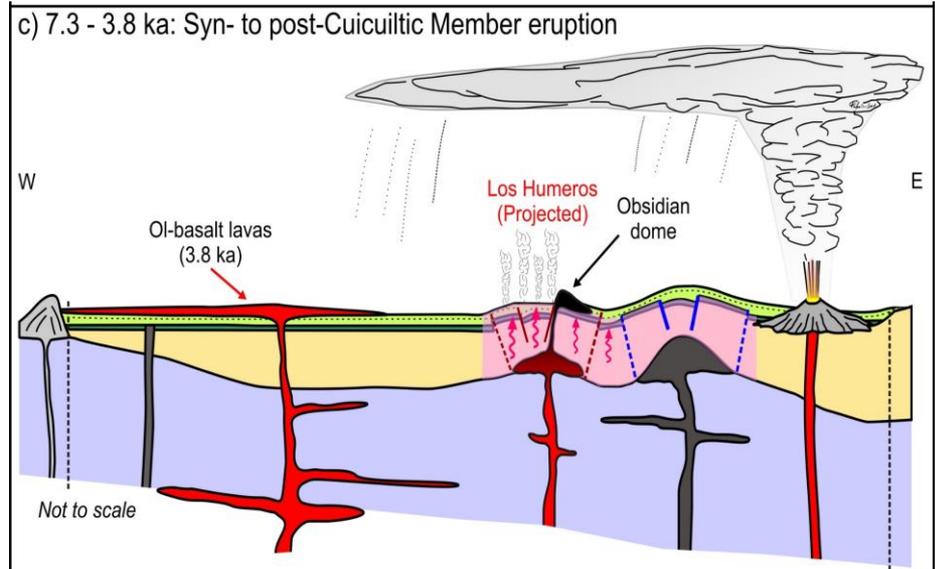
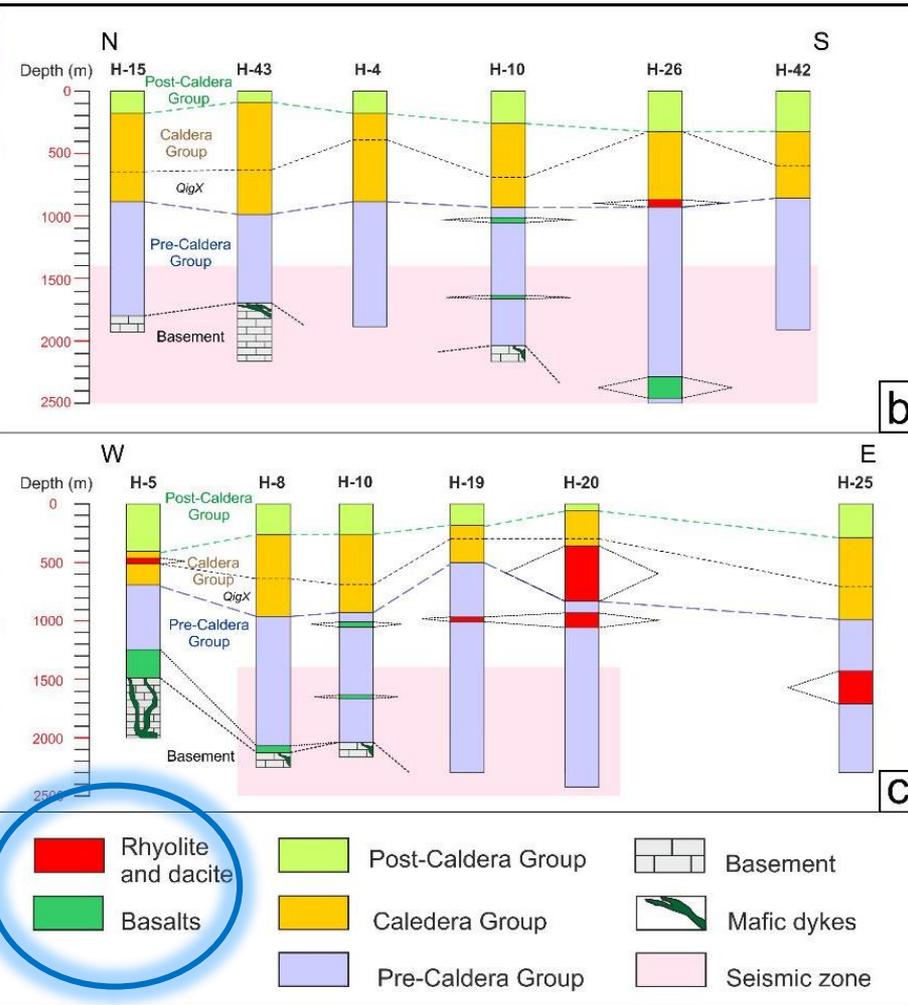
$$T_t = \frac{1}{2} L_g \times \frac{\sin(\theta + \alpha)}{\cos \theta}$$

e.g. Loma Blanca bulge depth of source of deformation:
 T = 425 ± 170 m.



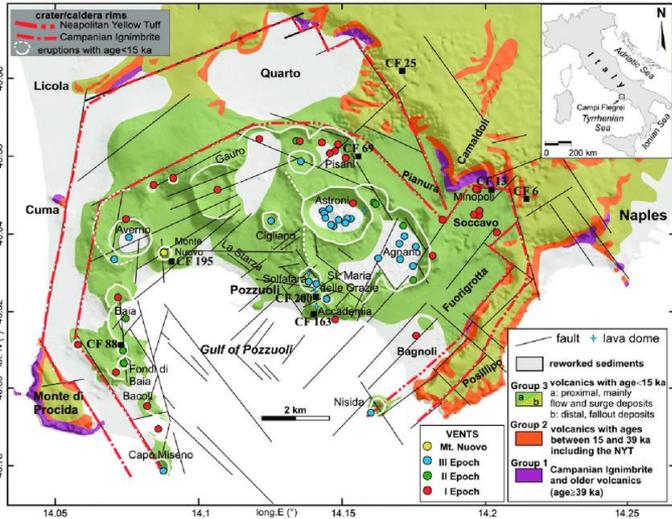
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Implications for subsurface setting

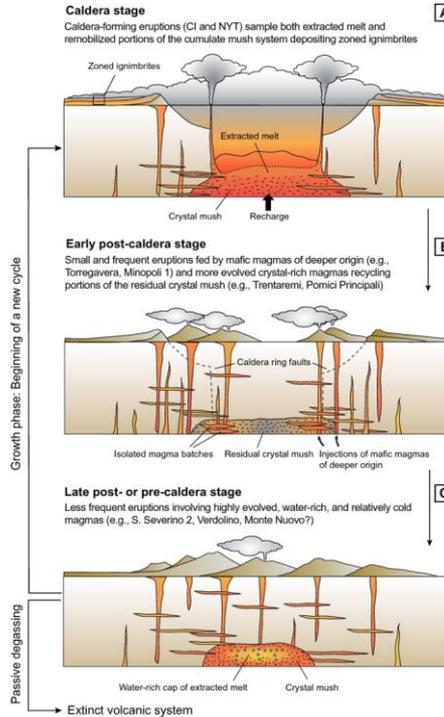


Urbani et al. (2020) Solid Earth

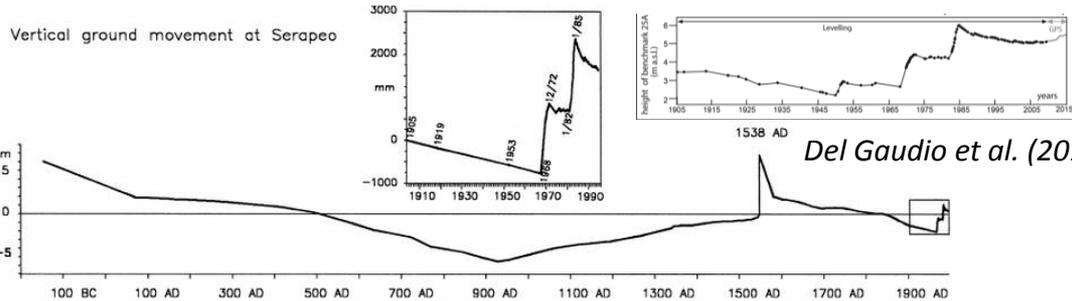
Analogue caldera – Campi Flegrei



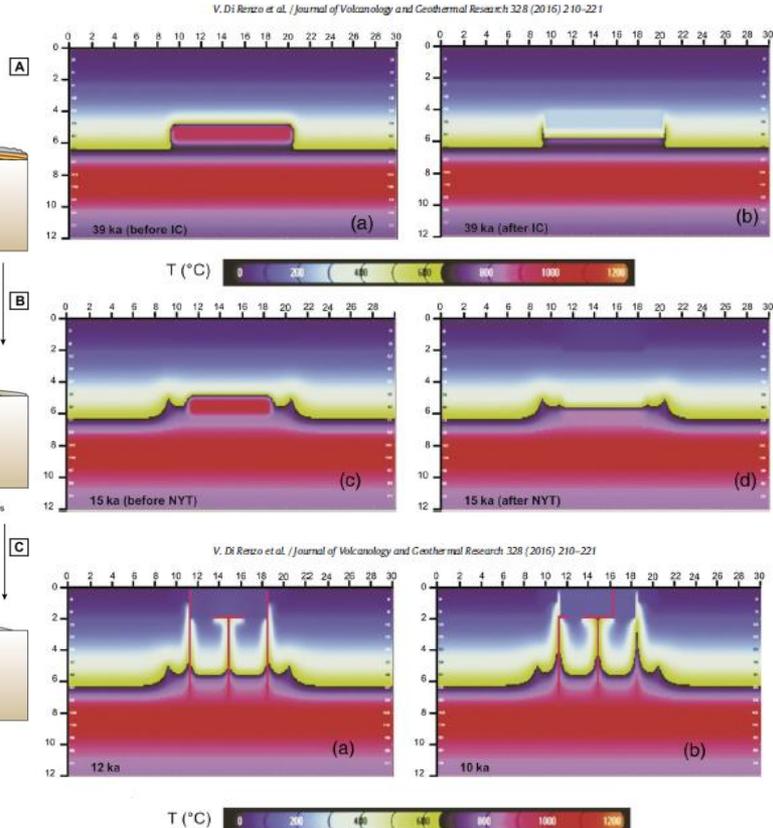
Isaia et al. (2016) GSAB



Forni et al. (2018) Sci Adv



Del Gaudio et al. (2010)



Di Renzo et al. (2016) JVGR



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- PETROLOGIC MODELING AT LH ALLOWED A STEP-CHANGE IN THE VOLCANOLOGIC CONCEPTUAL MODEL, FROM SINGLE MAGMA CHAMBER TO MULTILAYERED, TRANS-CRUSTAL RESERVOIRS
- UNPACKING AGE, VOLUME AND DEPTH OF THE MAGMATIC HEAT SOURCES IS THE FIRST ORDER NEED FOR GEOTHERMAL MODELING
- THERMAL MODELING OF HEAT CONDUCTION FROM THE MAGMA SOURCE MUST TAKE INTO ACCOUNT BOTH THE GEOMETRICAL CONSTRAINTS AND INCREMENTAL MASS BALANCES (RECHARGE vs WITHDRAWAL)
- THE CURRENT HEAT FLUX AT LH IS THE RESULT OF THE COMBINATION OF THE THERMAL RELAXATION OF THE XALTIPAN MAGMA CHAMBER + RECENT SHALLOWER INTRUSIONS + FRACTURE CONTROLLED FLUID ADVECTION
- MAGMATIC PRESSURE SOURCES THAT GENERATE BRITTLE DEFORMATION AND PERMEABILITY ARE POLYPHASED IN SPACE AND TIME
- DEVELOPMENT OF SHGS AT LH REQUIRES EFFORTS IN IMAGING THE VOLUME BETWEEN 7 AND 3 KM, WHERE MASS AND HEAT TRANSFER HAVE OCCURRED DURING THE LAST 10 KA

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Matteo Trolese (UniRoma Tre)
Paromita Deb (Aachen Uni)
CFE

Thomas Theye and Željka Žigovečki Gobac for assistance and suggestions during the EMPA and XRD analyses, Institut für Mineralogie und Kristallchemie, Universität Stuttgart, Stuttgart, Germany



This project has received funding from the European Union's Horizon 2020 research and innovation programme under grant agreement No. 727550

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